

WRPI Annual Conference, San Jose, April 6-7, 2017:

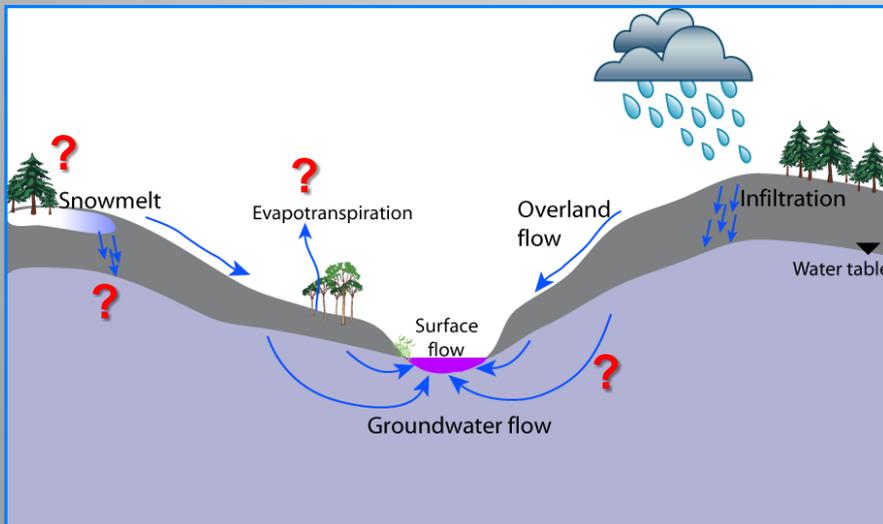
Climate Change Effects on Recharge in Headwater Catchments

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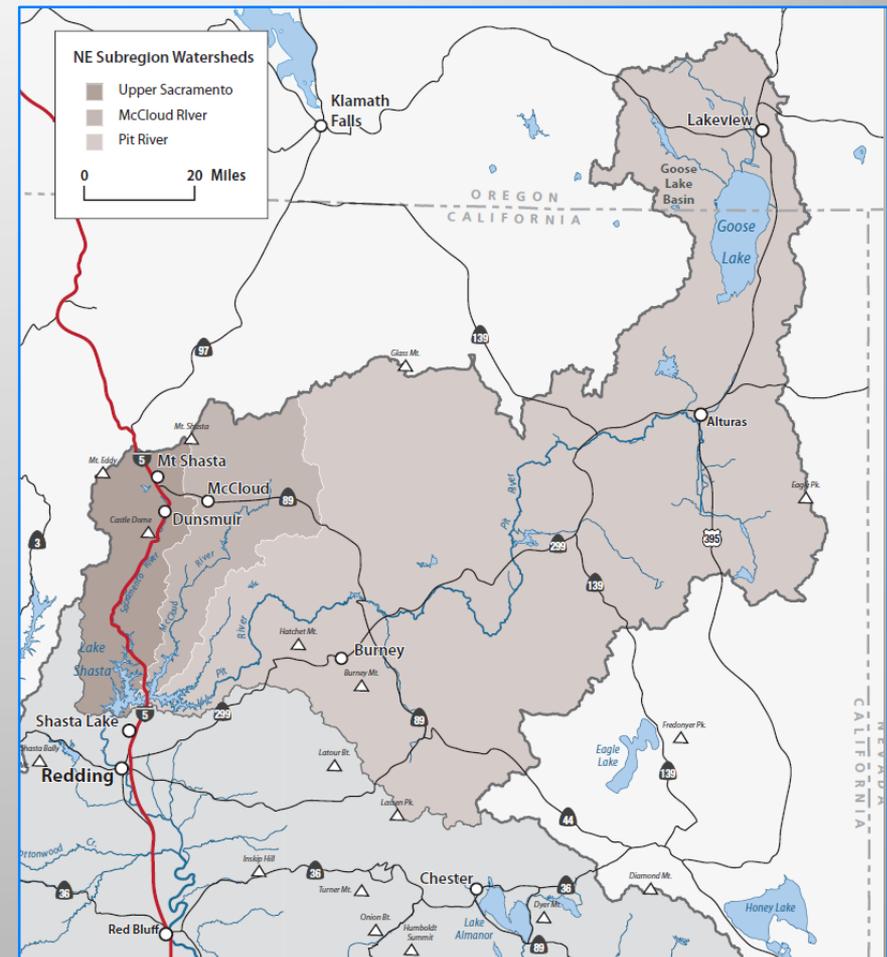
California AB 2480:

- “...it is state policy to recognize and define source watersheds as integral components of California's water system, and eligible for financing on an equivalent basis with other water infrastructure projects.”
- “...(funding for) projects with a demonstrated likelihood of increasing conditions for water and snow attraction, retention, and release under changing climate conditions.”



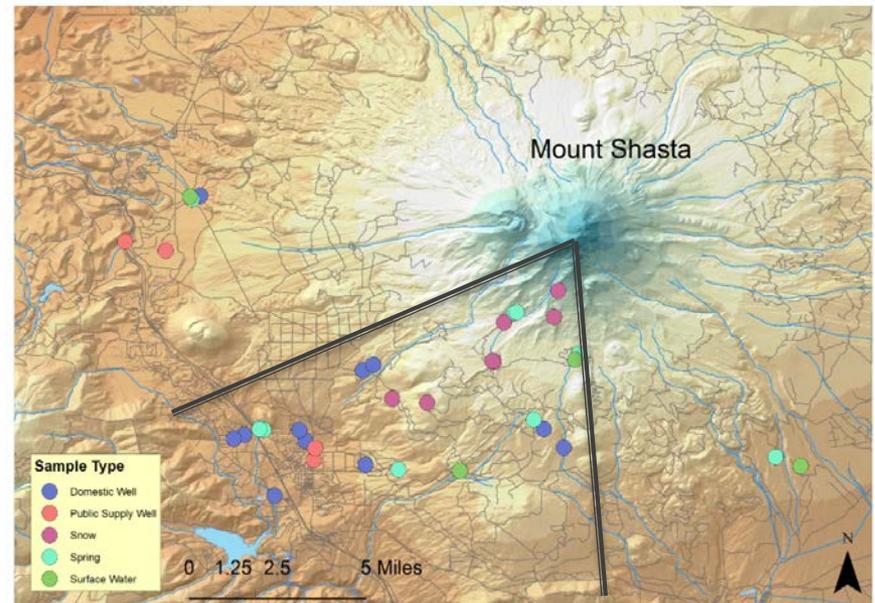
Volcanic, alpine hydrologic setting

- Highly permeable surface materials
- Stream generation from springs
 - Upper Sacramento, McCloud, and Pit River headwaters
 - Springs and wells provide supply for logging operations, towns of Mount Shasta and Weed, bottled water facility, etc.
- ET likely energy (rather than water) limited

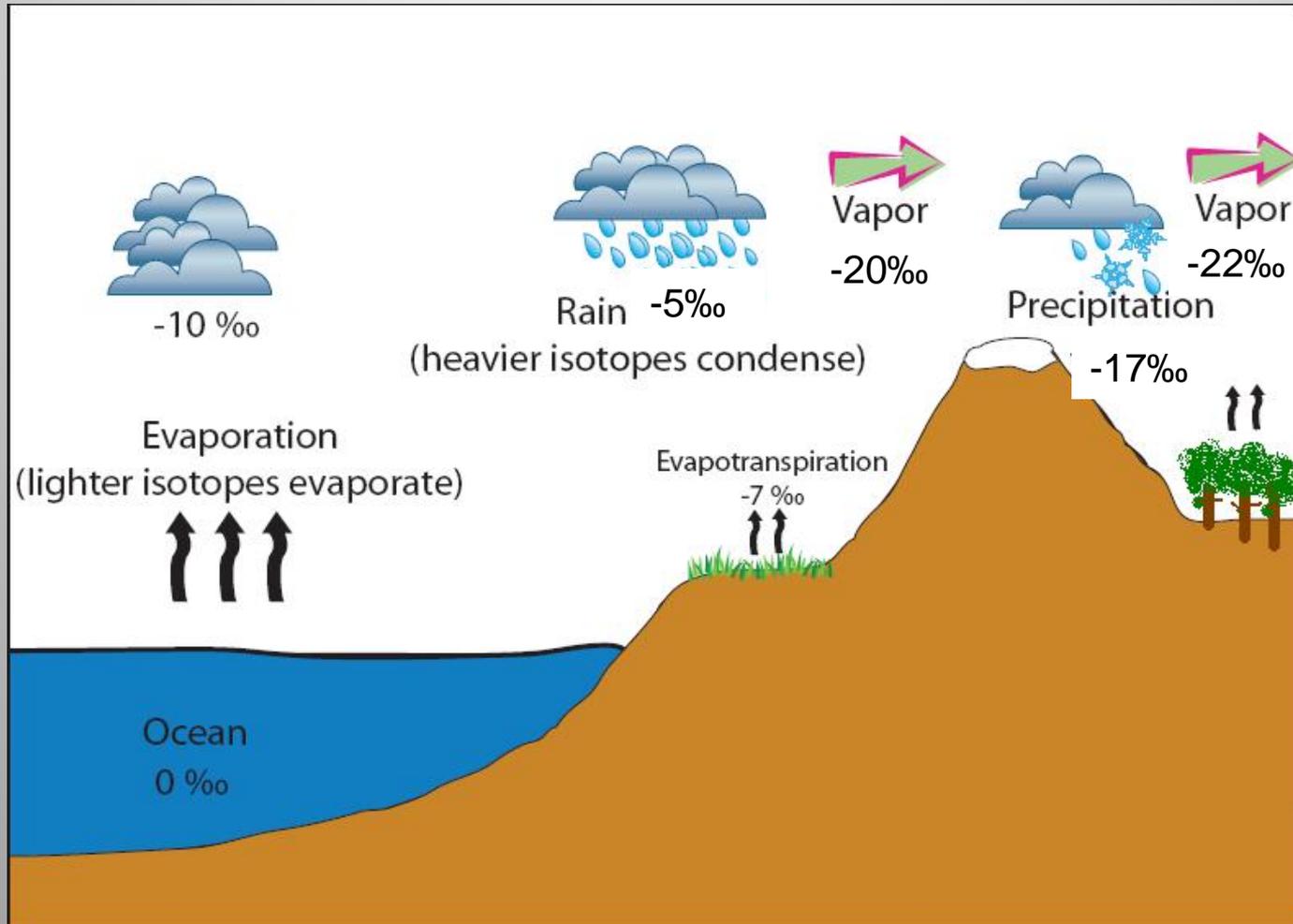


Sampling in headwater areas

- Limited sampling locations
- Sampled during extreme drought (May and Sept. of 2015)



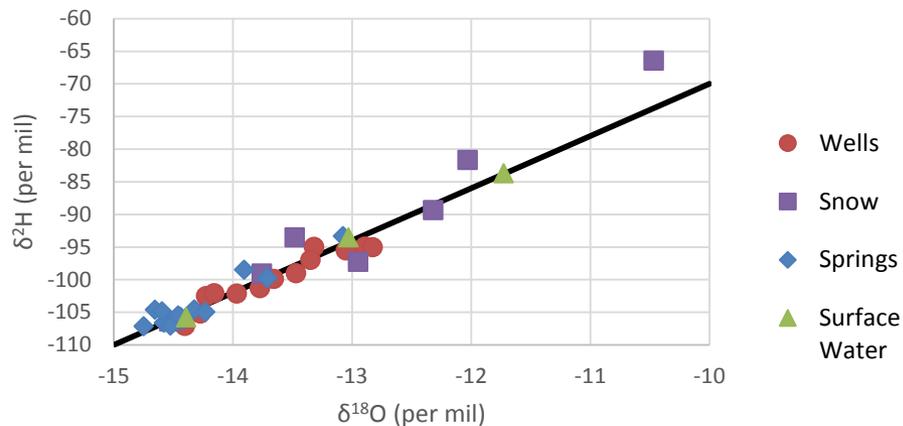
Stable isotope signature: How it works



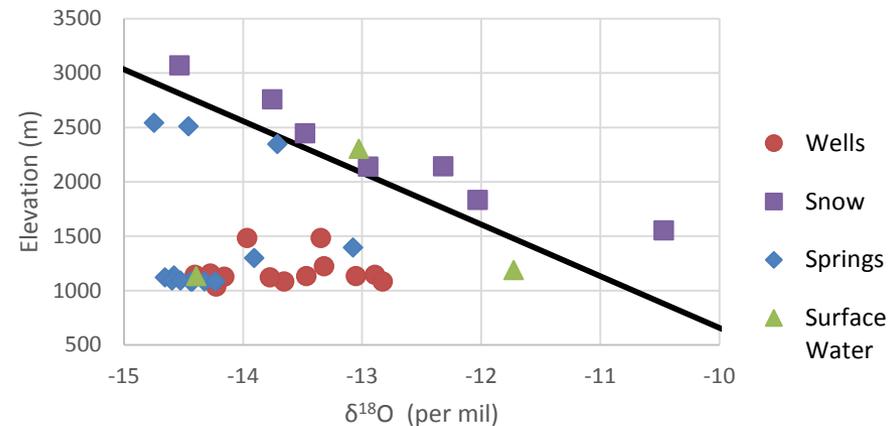
Stable isotopes: recharge elevation

- All samples fall on or near the Global Meteoric Water Line
- No $\delta^{18}\text{O}$ results $< -15\text{‰}$
- Precipitation (snow) follows 'lapse rate' (from Rose et al., 1996)
- Points that fall below the line indicate a source area at higher elevation

$\delta^{18}\text{O}$ vs. $\delta^2\text{H}$

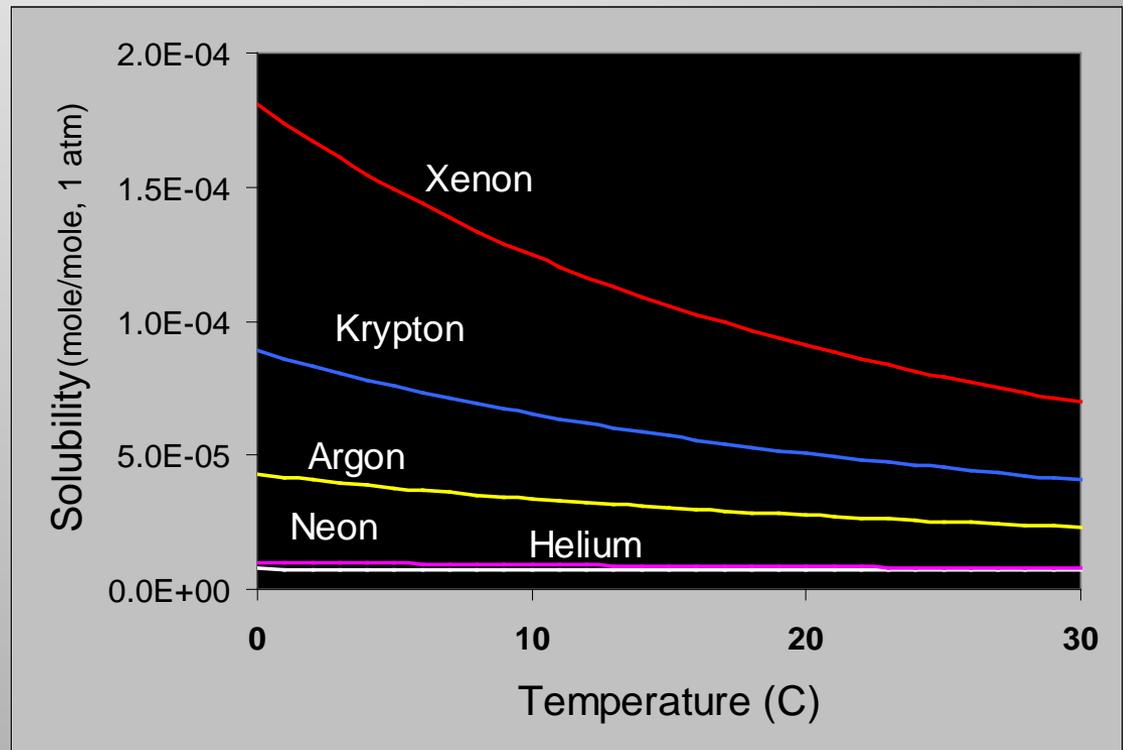
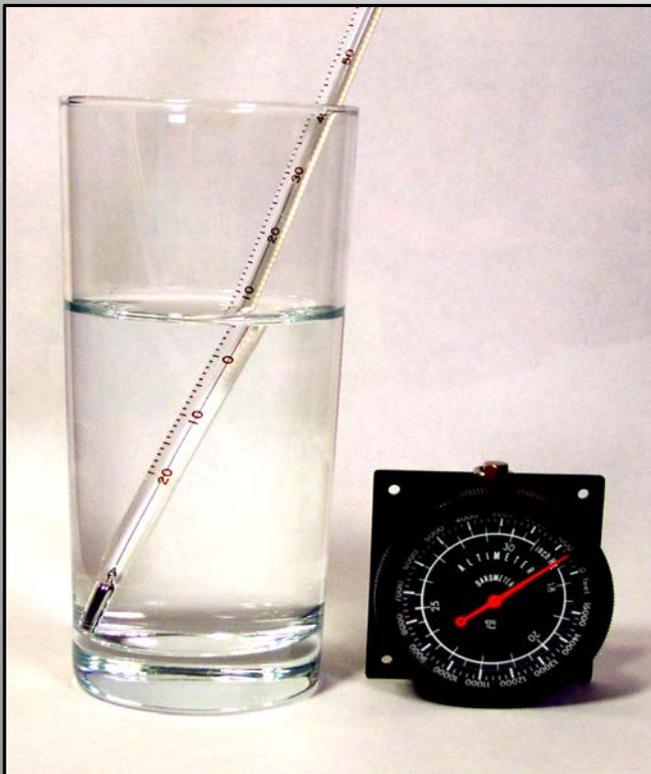


$\delta^{18}\text{O}$ vs. Sample Elevation

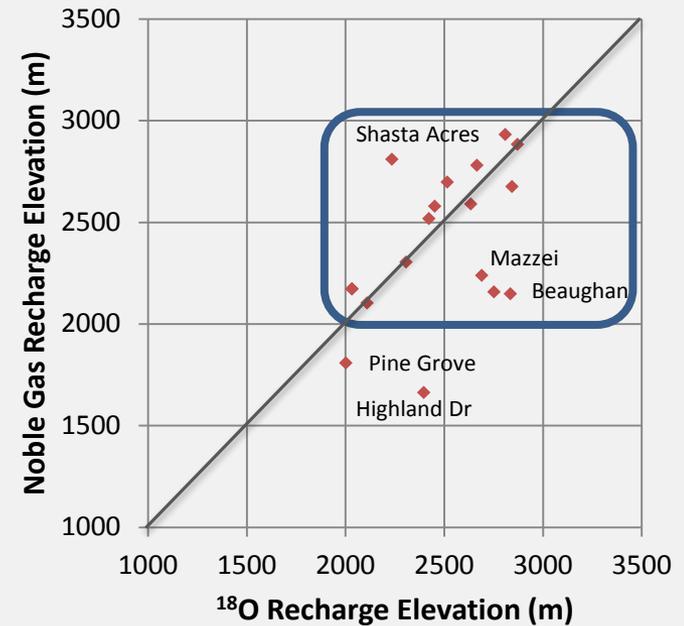
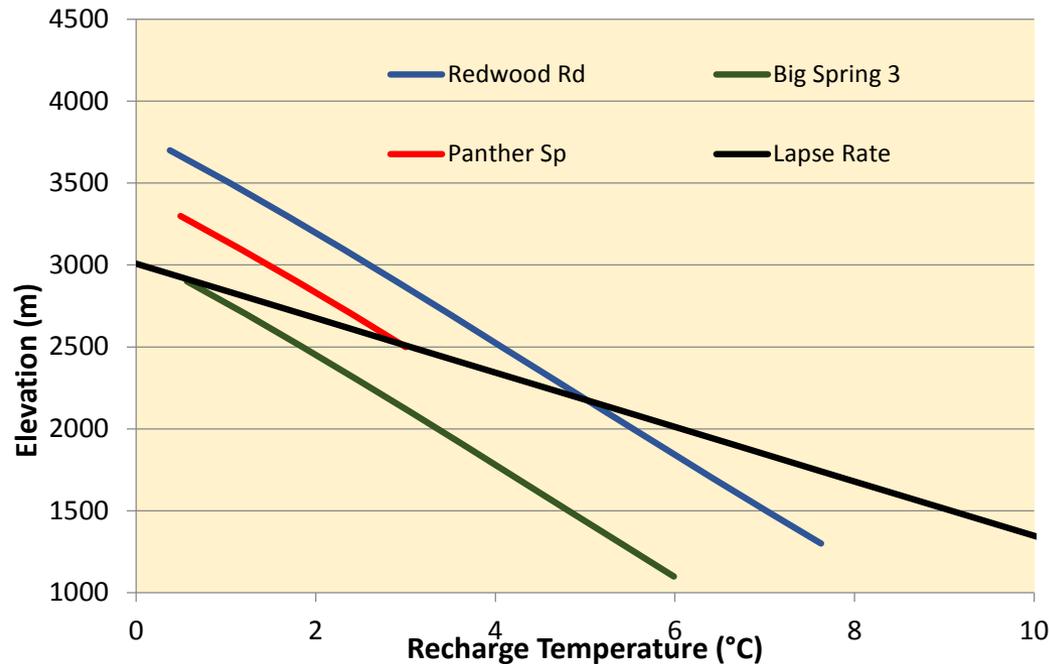


Noble gas signature: How it works

- Solubility depends on temperature and pressure
- Heavier gases have stronger T dependence



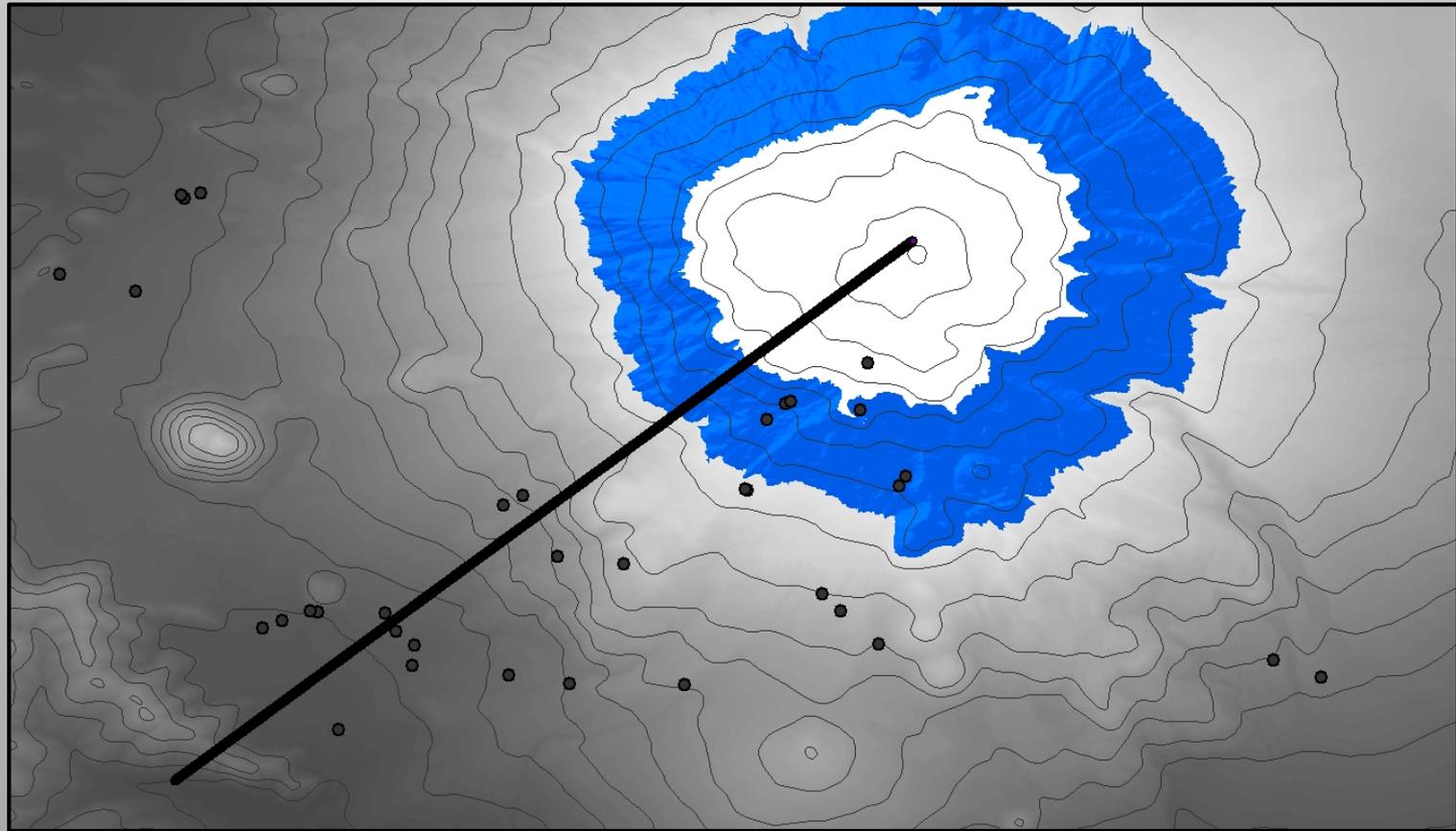
Noble gases: recharge elevation



- 3 examples
- Constraints:
 - elevation (top of mountain to sampling elevation)
 - temperature (greater than 0°C, less than discharge temperature)

Good agreement between independent methods

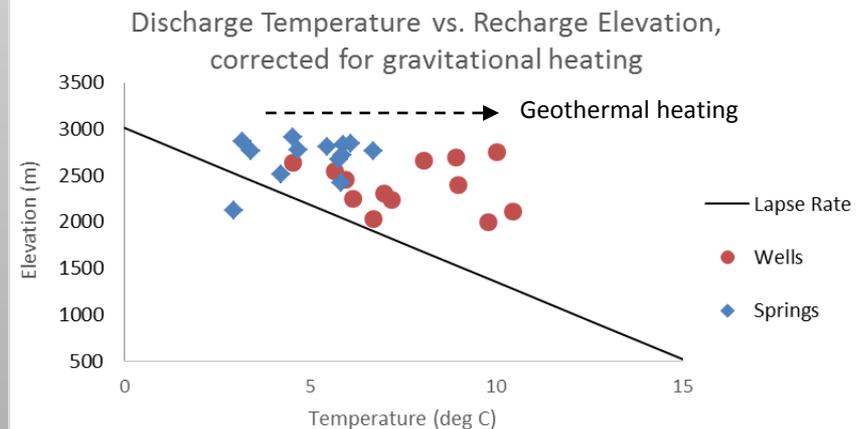
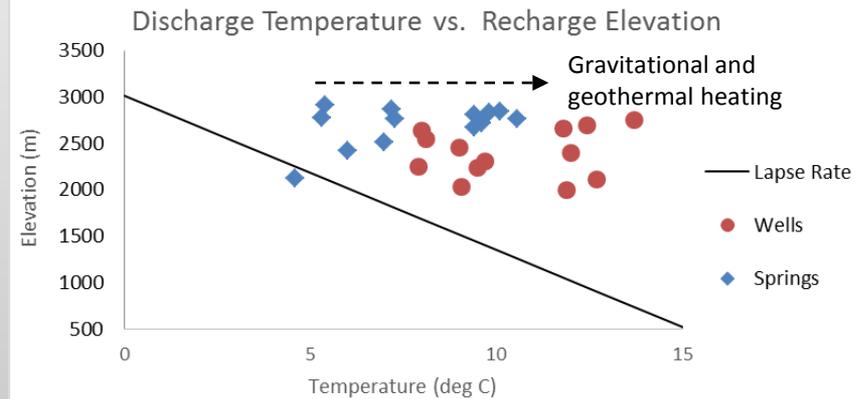
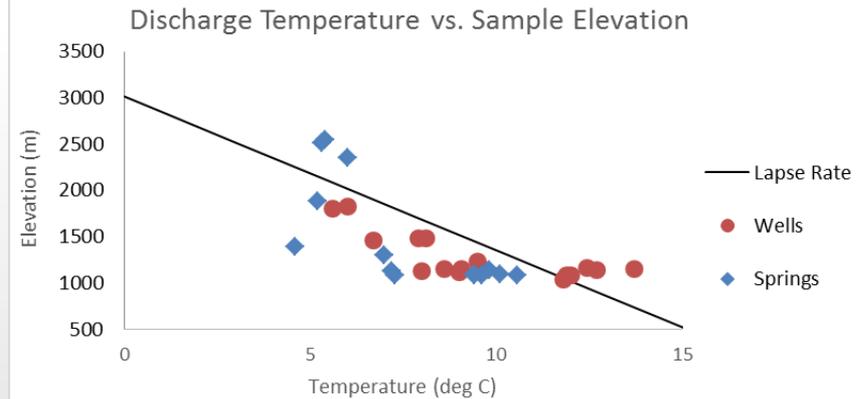
Delineating the recharge area



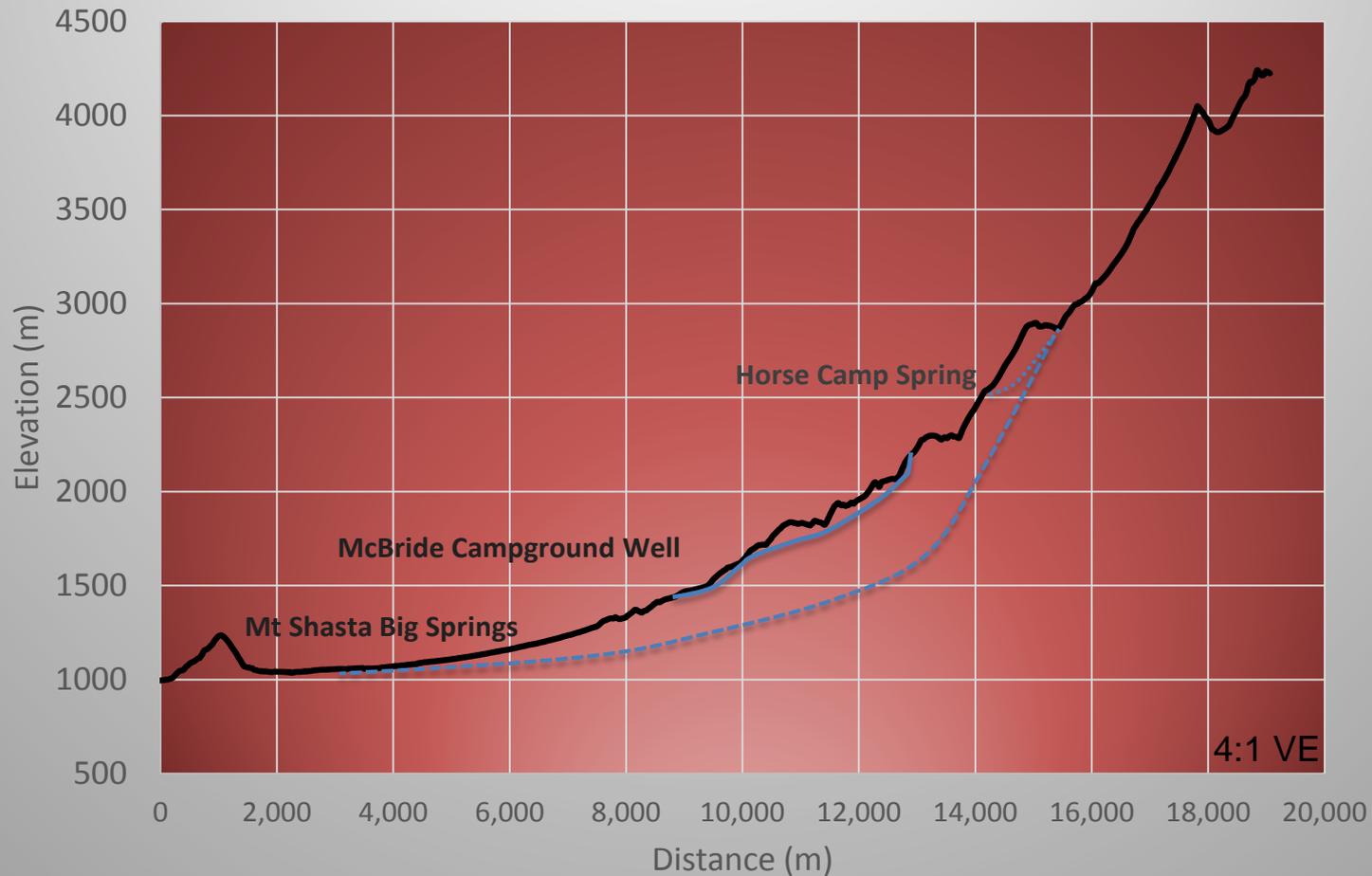
- Elevation range 2100-2900 m
- Do not observe signals from higher or lower elevations

Subsurface Heating

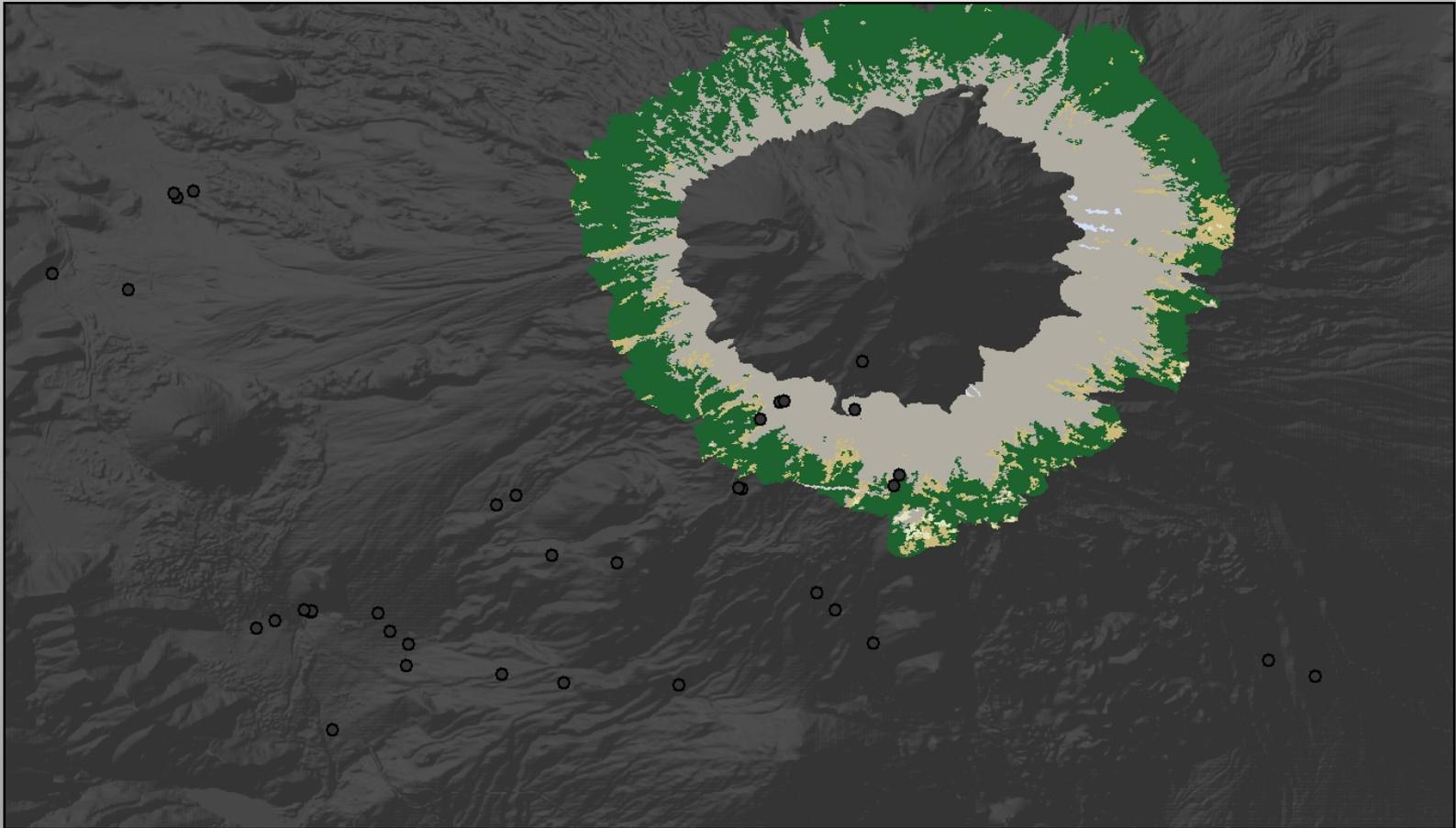
- The rate of change of thermal energy in a parcel of groundwater is the sum of (Manga and Kirchner, 2004):
 - gravitational potential energy dissipation
 - heat transfer to/from the surface by circulating water via conduction (negligible in this setting)
 - geothermal heating
- For a 8°C DT-RT difference and a geothermal gradient of $15^{\circ}\text{C}/\text{km}$, a maximum flow depth of appx. 500 m is calculated



Delineating groundwater flow on Mount Shasta



Land cover in the recharge area



Forest ET likely plays a key role in limiting percolation and recharge

There may be a 'sweet spot' in tree coverage

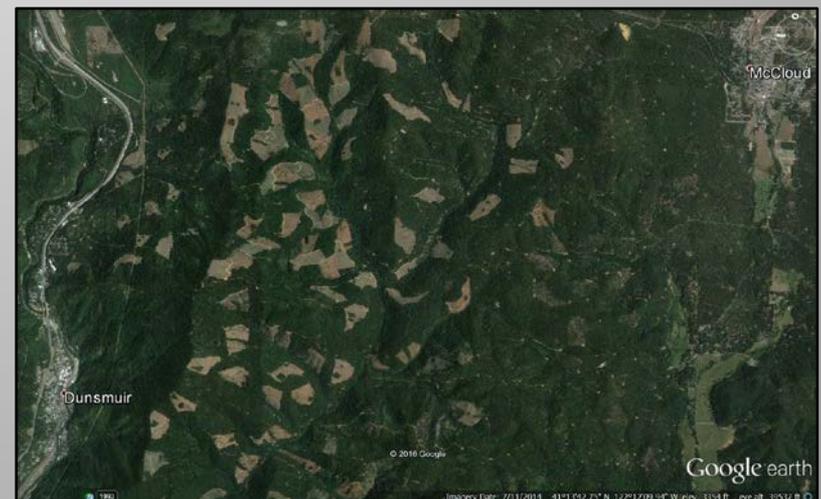
- Too many trees

- Increased ET
- Decreased groundwater flow to streams
- Insufficient openings for snow accumulation



- Too few trees

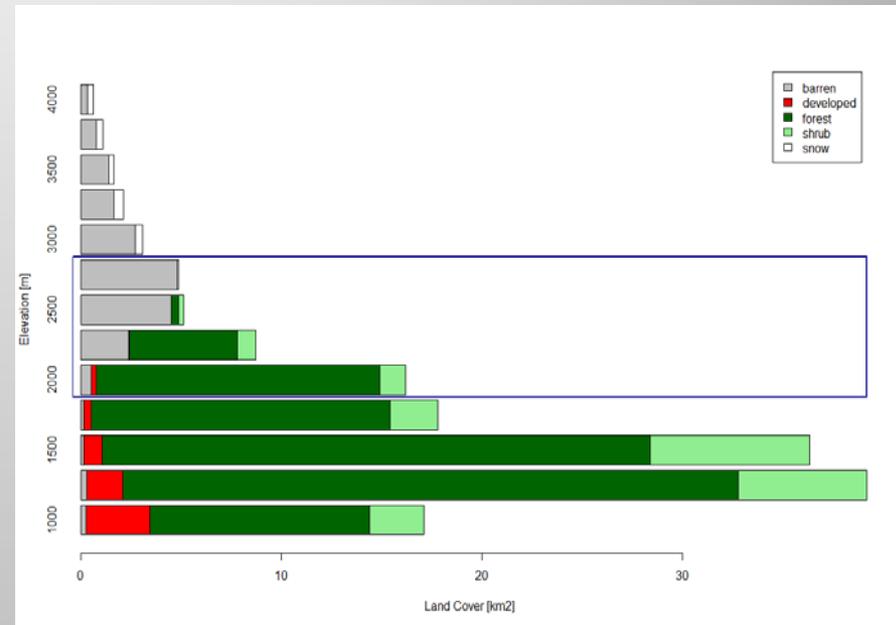
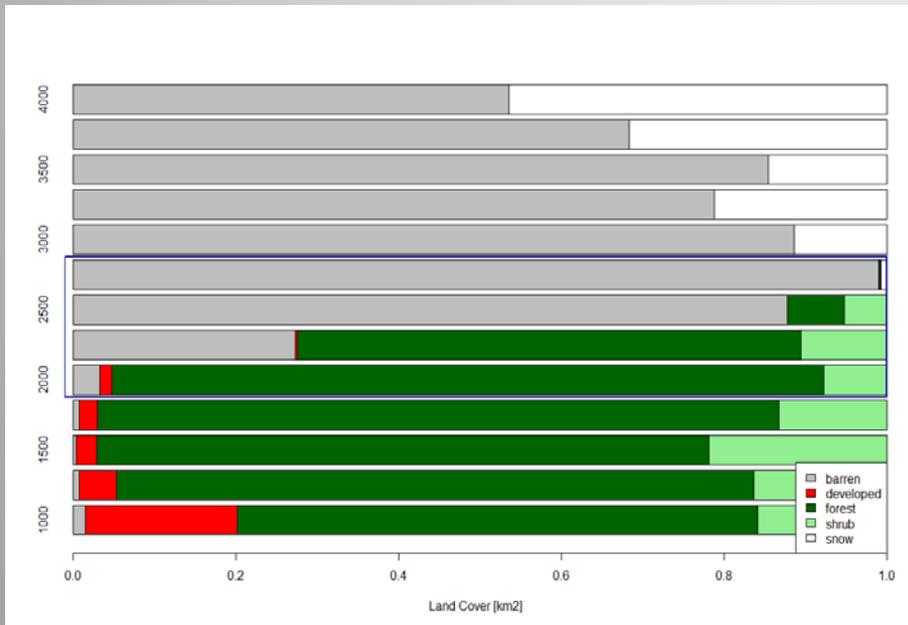
- Increased runoff
- Increased sediment
- Increased stream temperature
- Not enough shade to extend snowmelt season



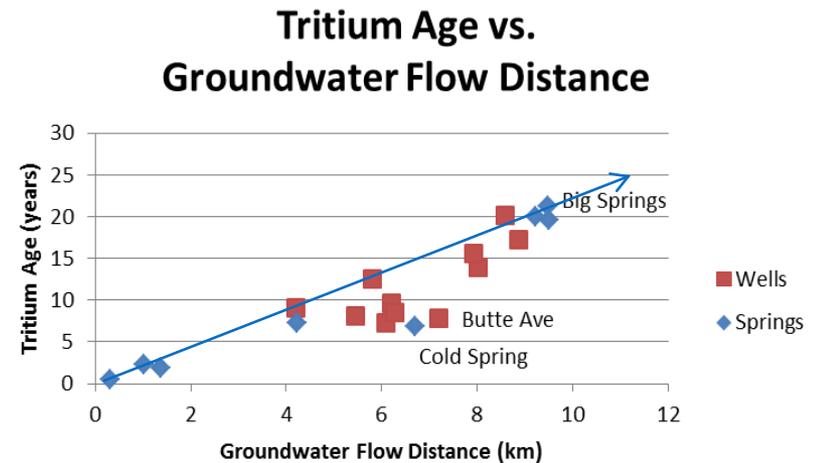
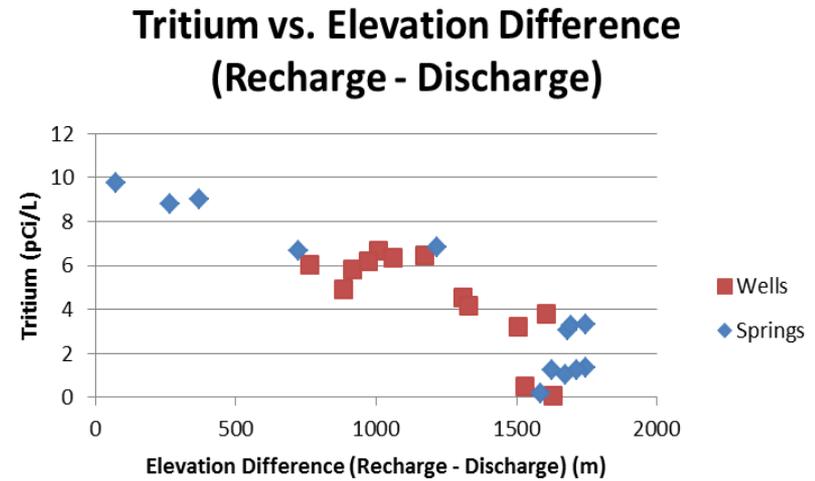
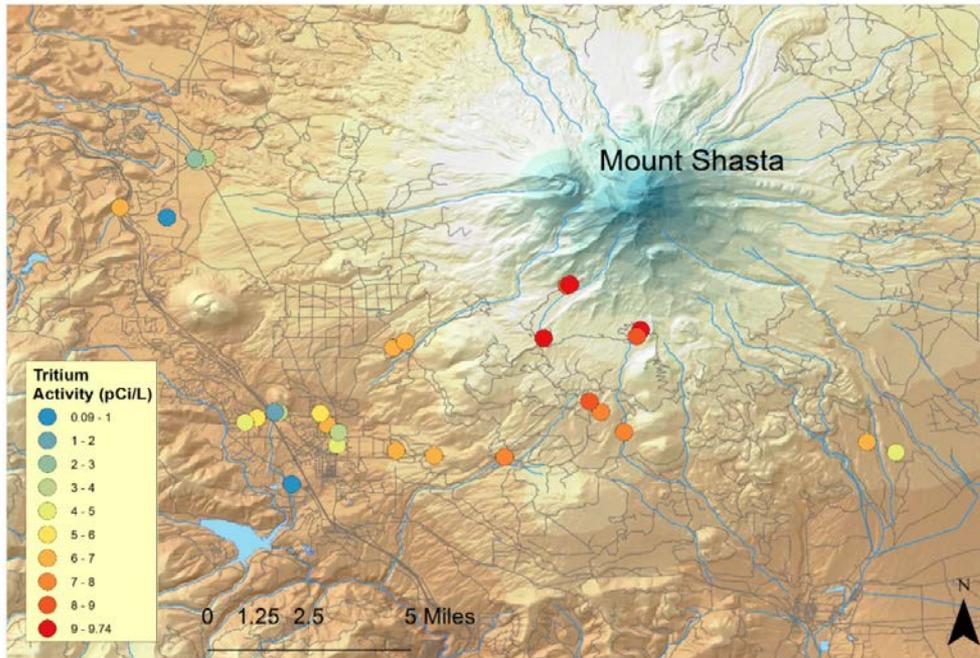
Conclusions

- Headwater areas provide critical late season flow and ecosystem functions
- On Mount Shasta, stable isotopes and noble gas recharge temperatures indicate that recharge occurs predominantly over the elevation range 2200-2900 m (7200 ft to 9500 ft)
- Higher elevations are disproportionately represented because of high precipitation rates and low evapotranspiration over bare ground
- The warmer future will bring a higher treeline, smaller area of bare rock, and likely lower recharge rates.

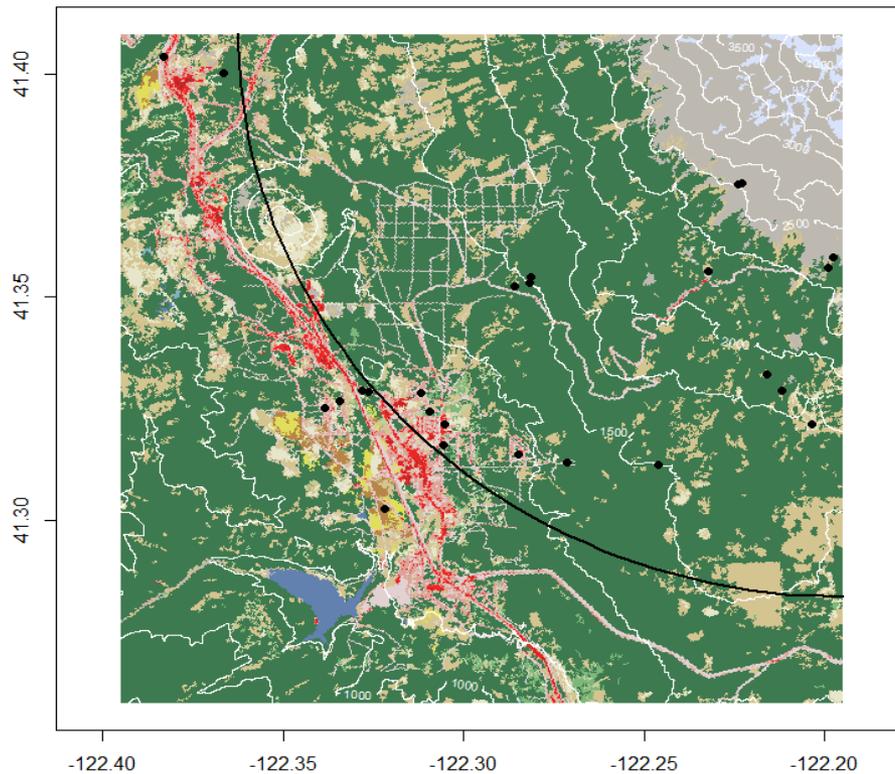
Relationships between elevation and land cover/area



Groundwater ages and flow rates



Land coverage, ET, and recharge



- How much of the melting snow goes to ET instead of infiltrating?
- Changes due to warming climate?

Effects of climate change on ET and groundwater recharge

